



ILMATIETEEN LAITOS
METEOROLOGISKA INSTITUTET
FINNISH METEOROLOGICAL INSTITUTE

Earth System modelling School 2009

3-7 August 2009, Helsinki, Finland

Atmosphere in the Earth system

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Outline ::

- **Role of the atmosphere in the Earth system**
 - 1. Interaction with the surrounding Universe**
 - Upper boundary condition of the Earth system
 - 2. Interaction with the underlying surface**
 - Lower boundary condition of the atmosphere
 - Non-linear interaction with other Earth system components
 - 3. Internal dynamics**



1. Interaction with the surrounding Universe

- **Solar wind of high energy particles + cosmic rays**
- **Lunar interaction through gravitation (ocean tides)**
- **Solar irradiance**



1. Interaction with the surrounding Universe

- Solar wind of high energy particles + cosmic rays
- Lunar interaction through gravitation (ocean tides)
- **Solar irradiance**

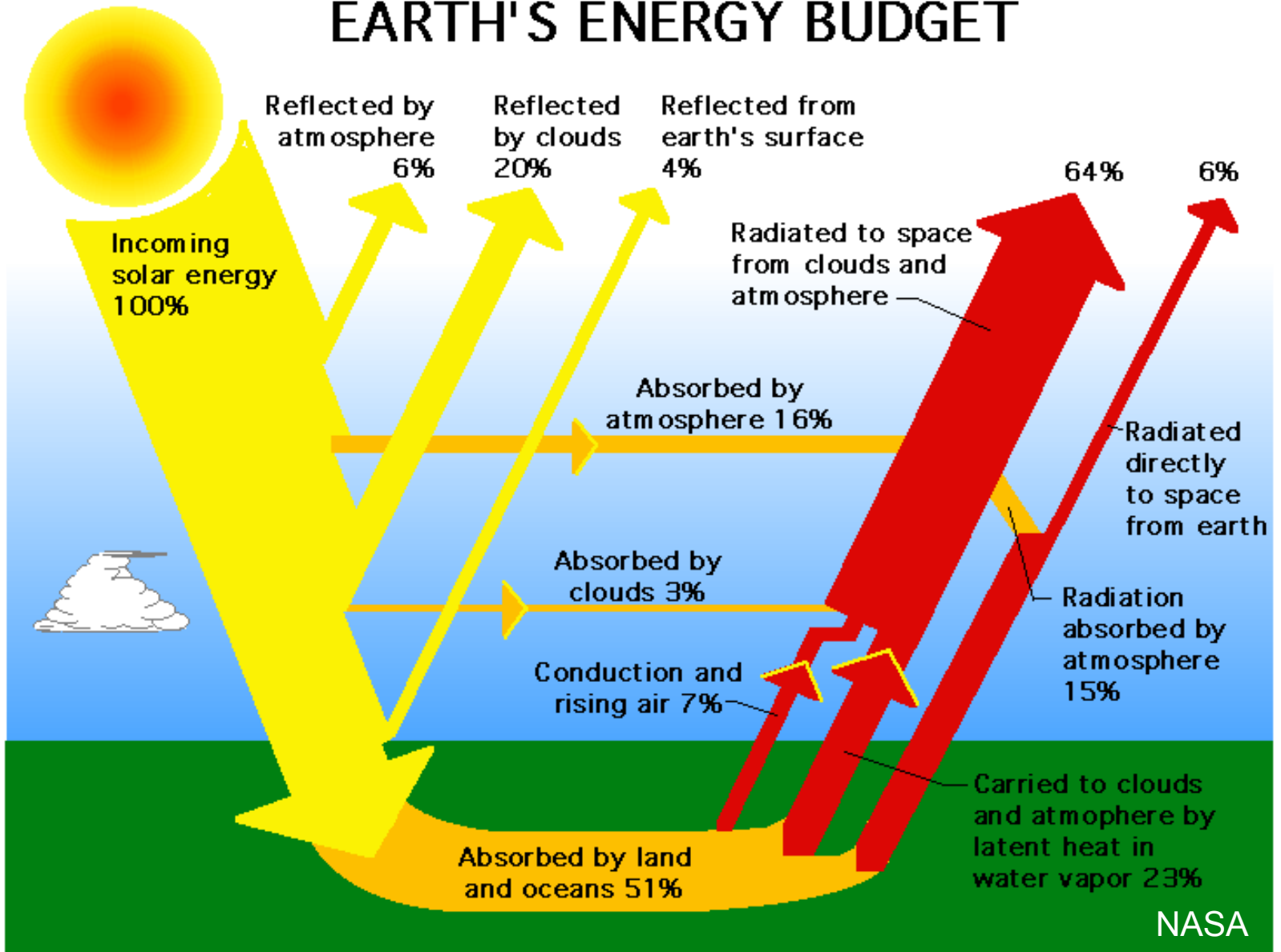


Radiation budget of the Earth

- **incoming (shortwave) radiation from the Sun**
- **reflected (shortwave) radiation from the Earth**
- **outgoing thermal (longwave) from the Earth**



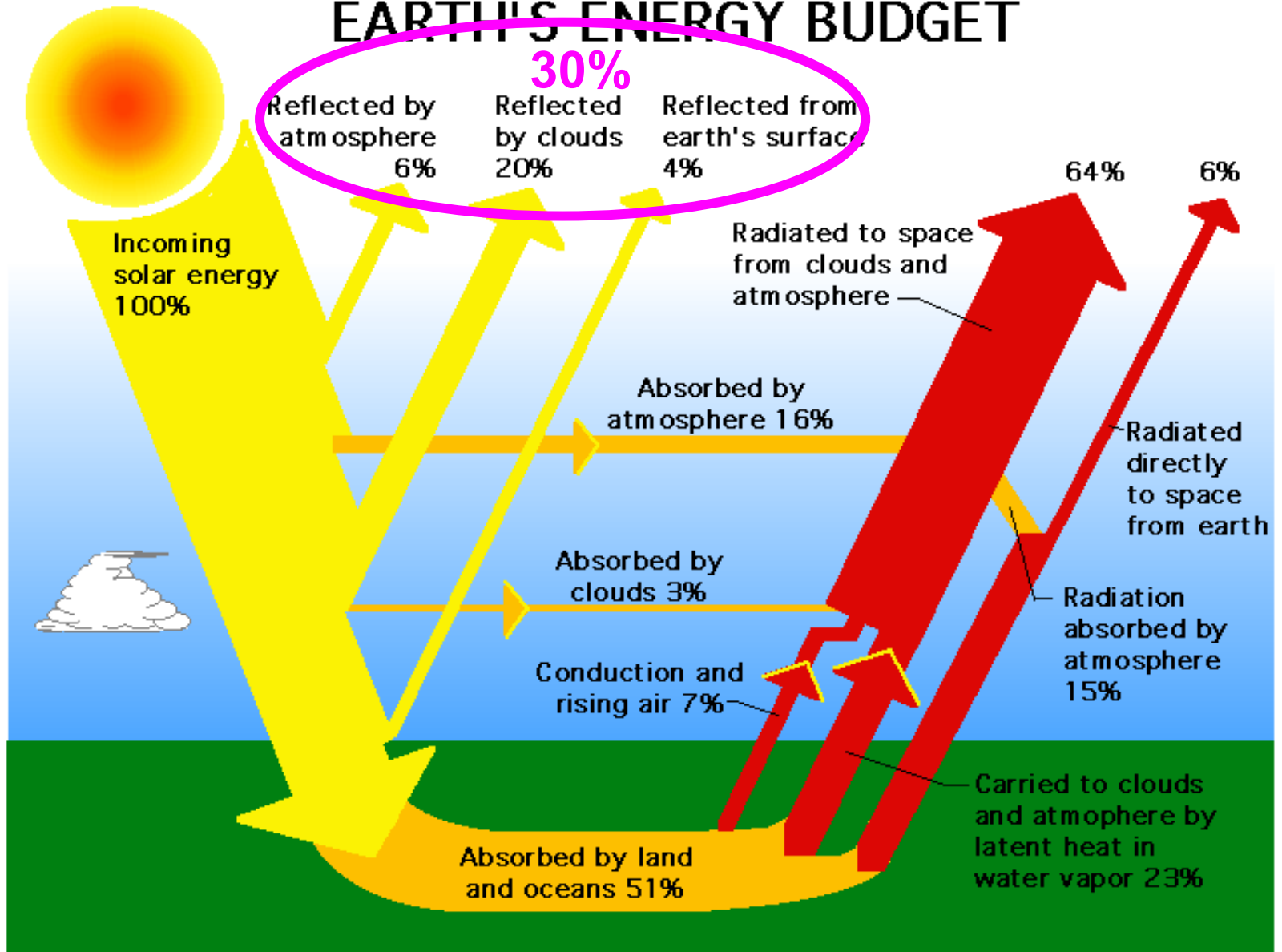
EARTH'S ENERGY BUDGET



NASA

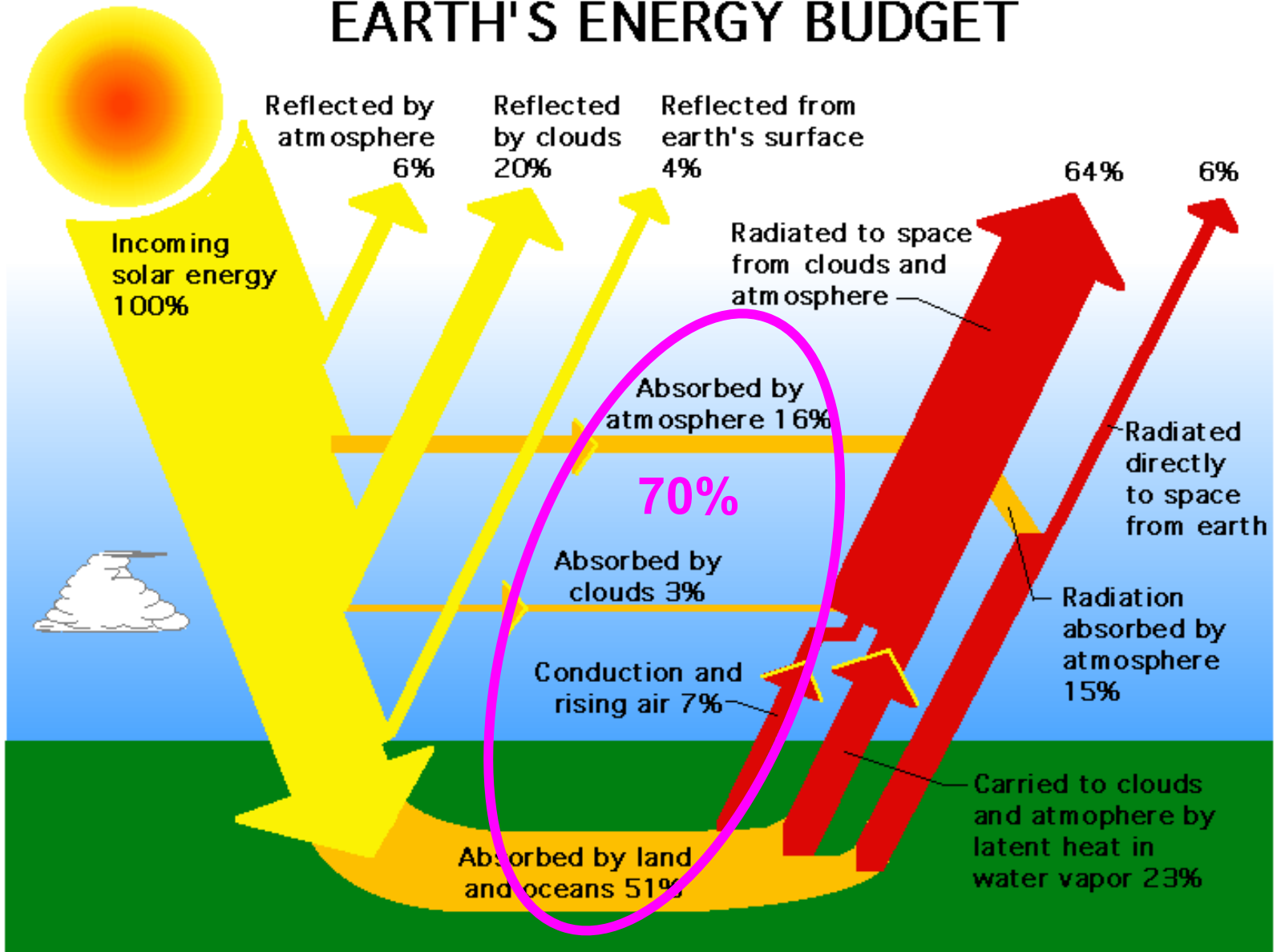


EARTH'S ENERGY BUDGET



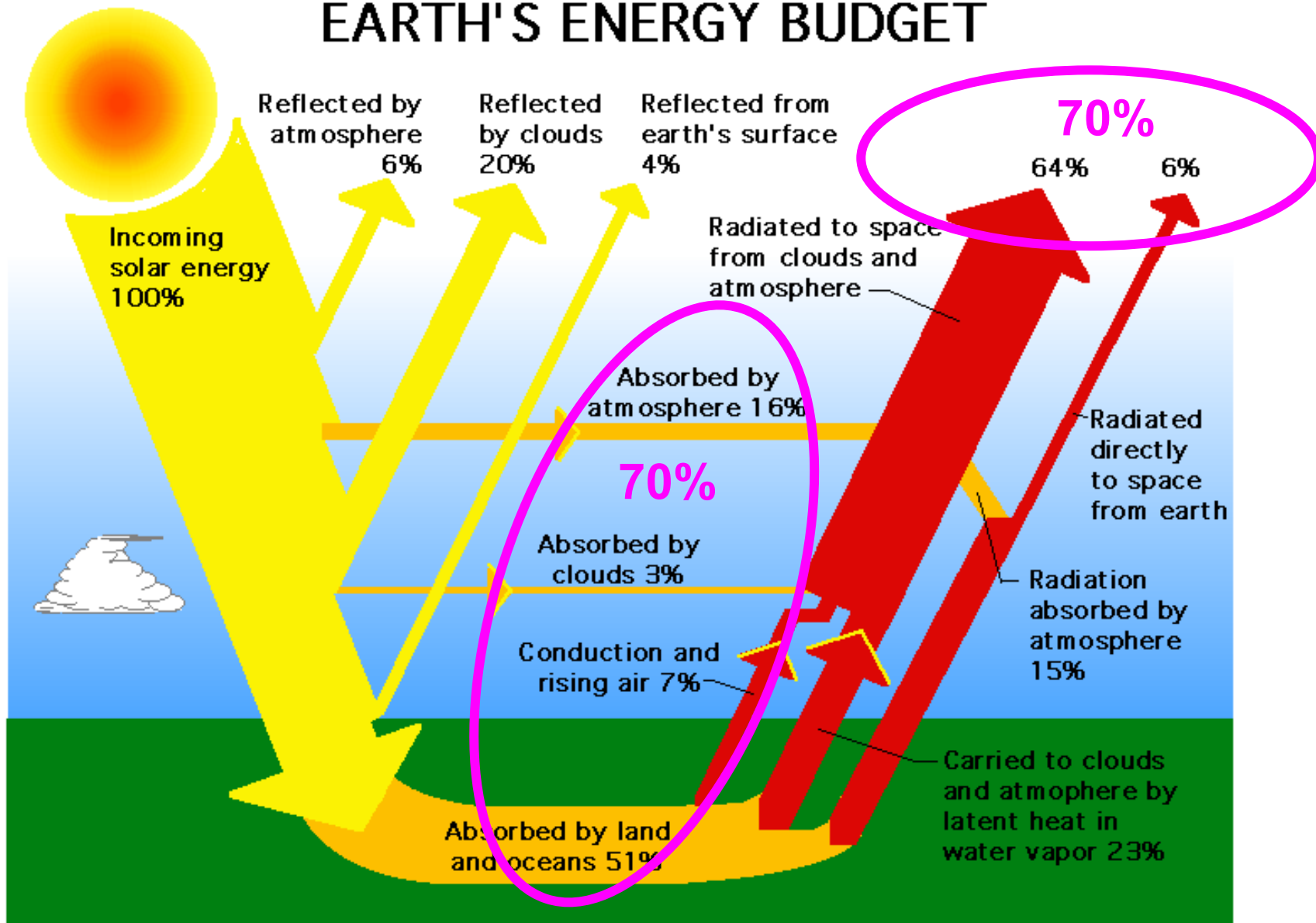


EARTH'S ENERGY BUDGET



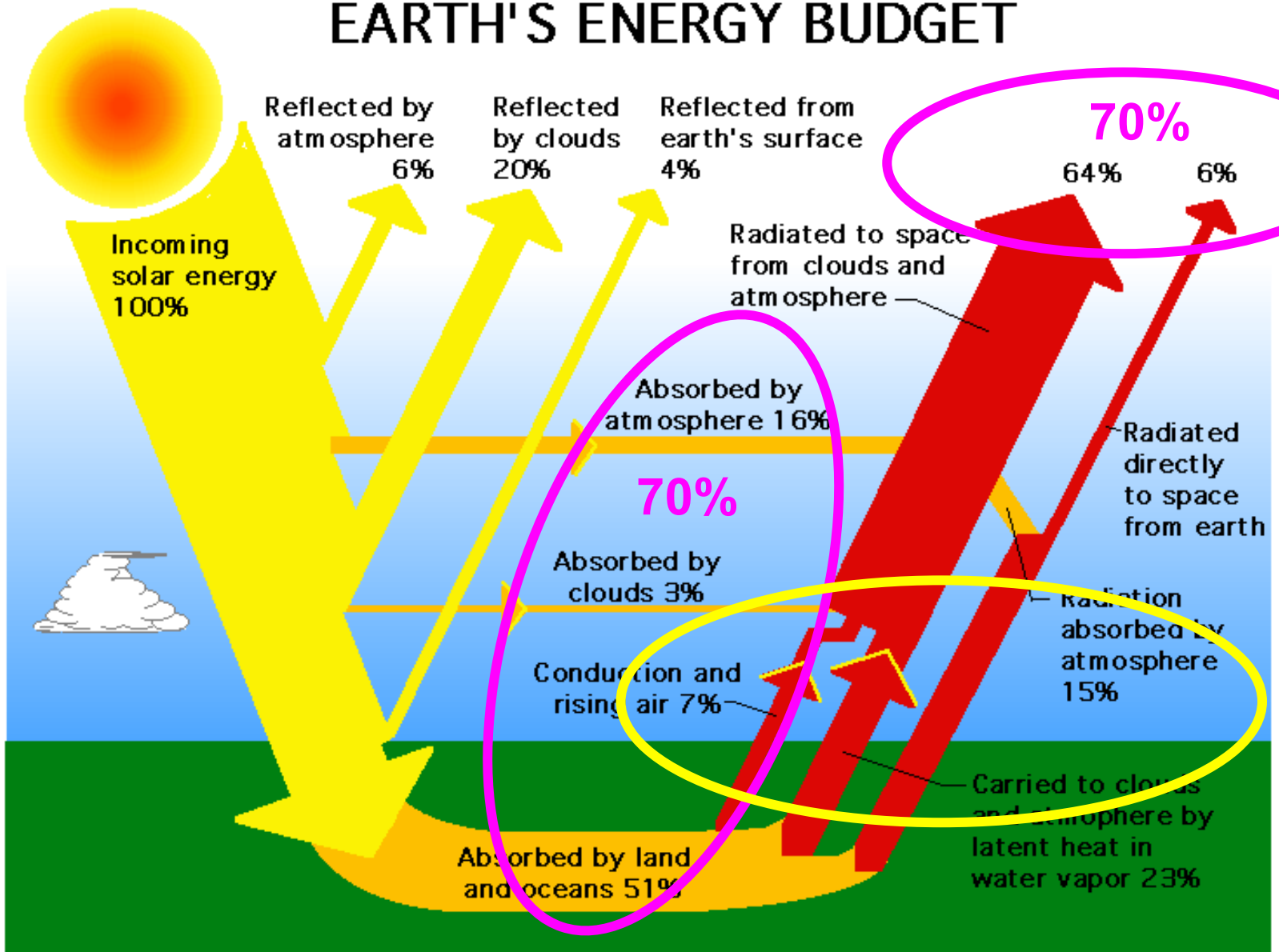


EARTH'S ENERGY BUDGET



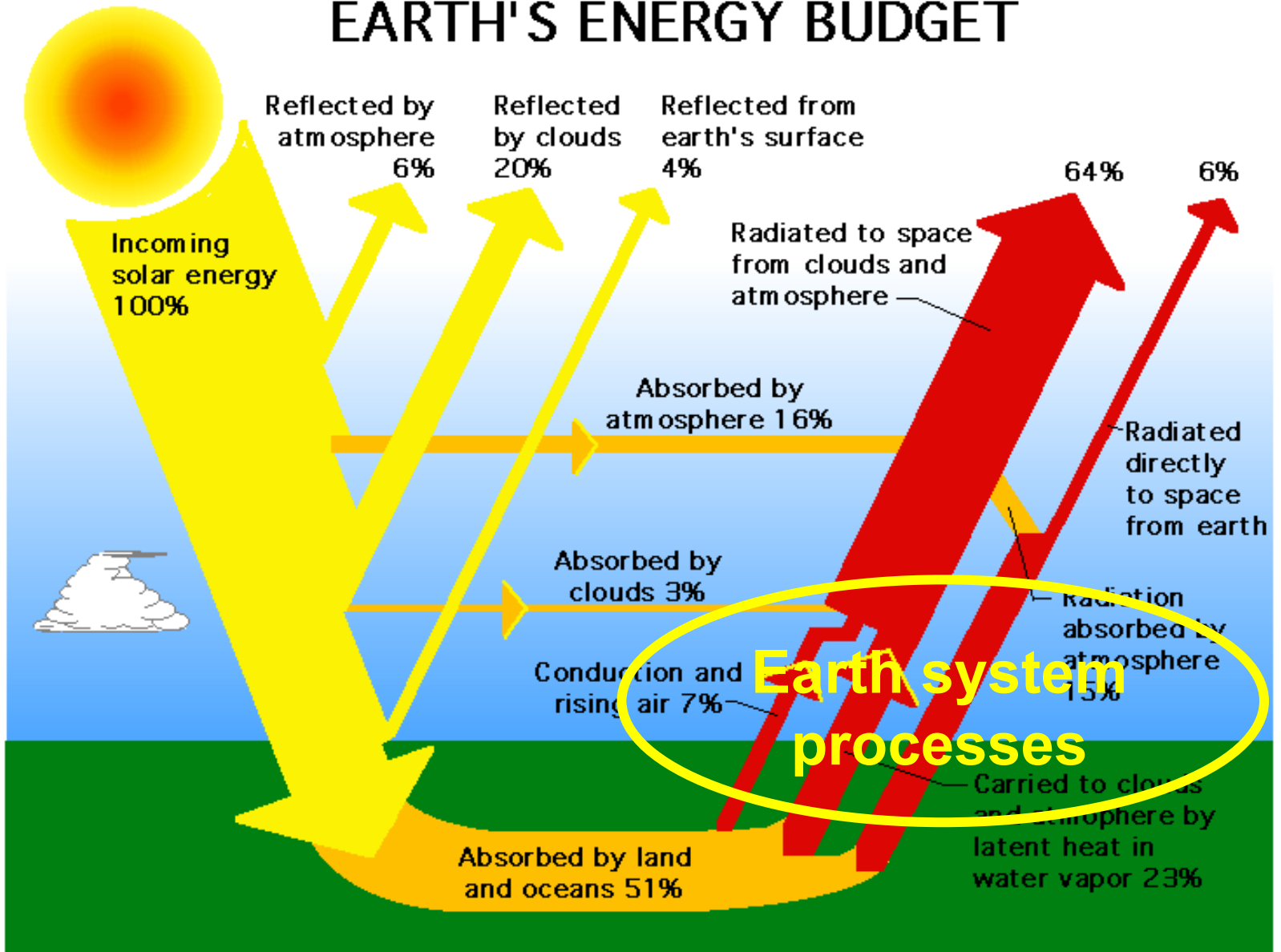


EARTH'S ENERGY BUDGET





EARTH'S ENERGY BUDGET





Upper boundary condition

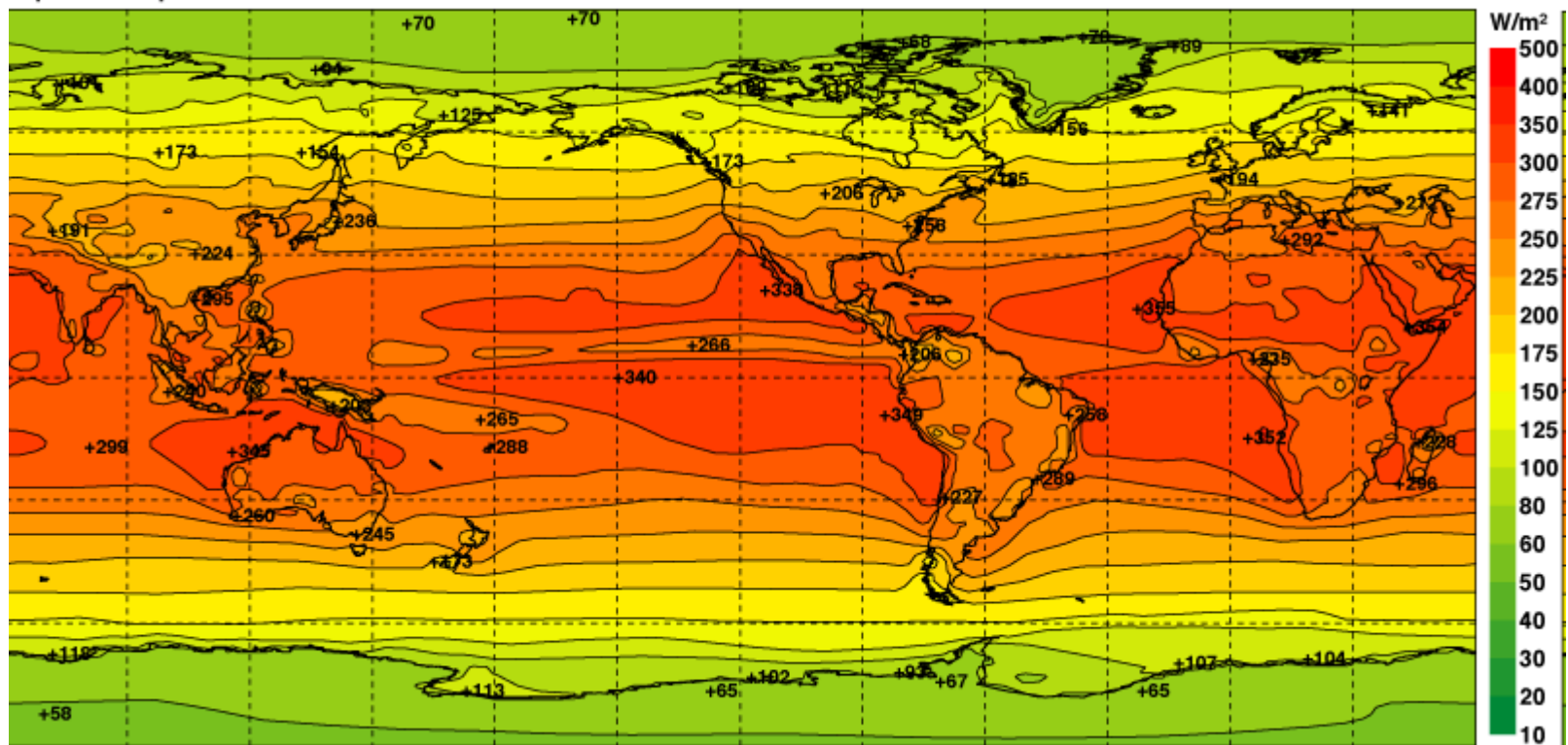
- **purely radiative**
- **Next: net shortwave and longwave radiation (distribution at the top of the atmosphere)**



Incoming net solar radiation (absorbed)

Top of atmosphere net solar radiation.

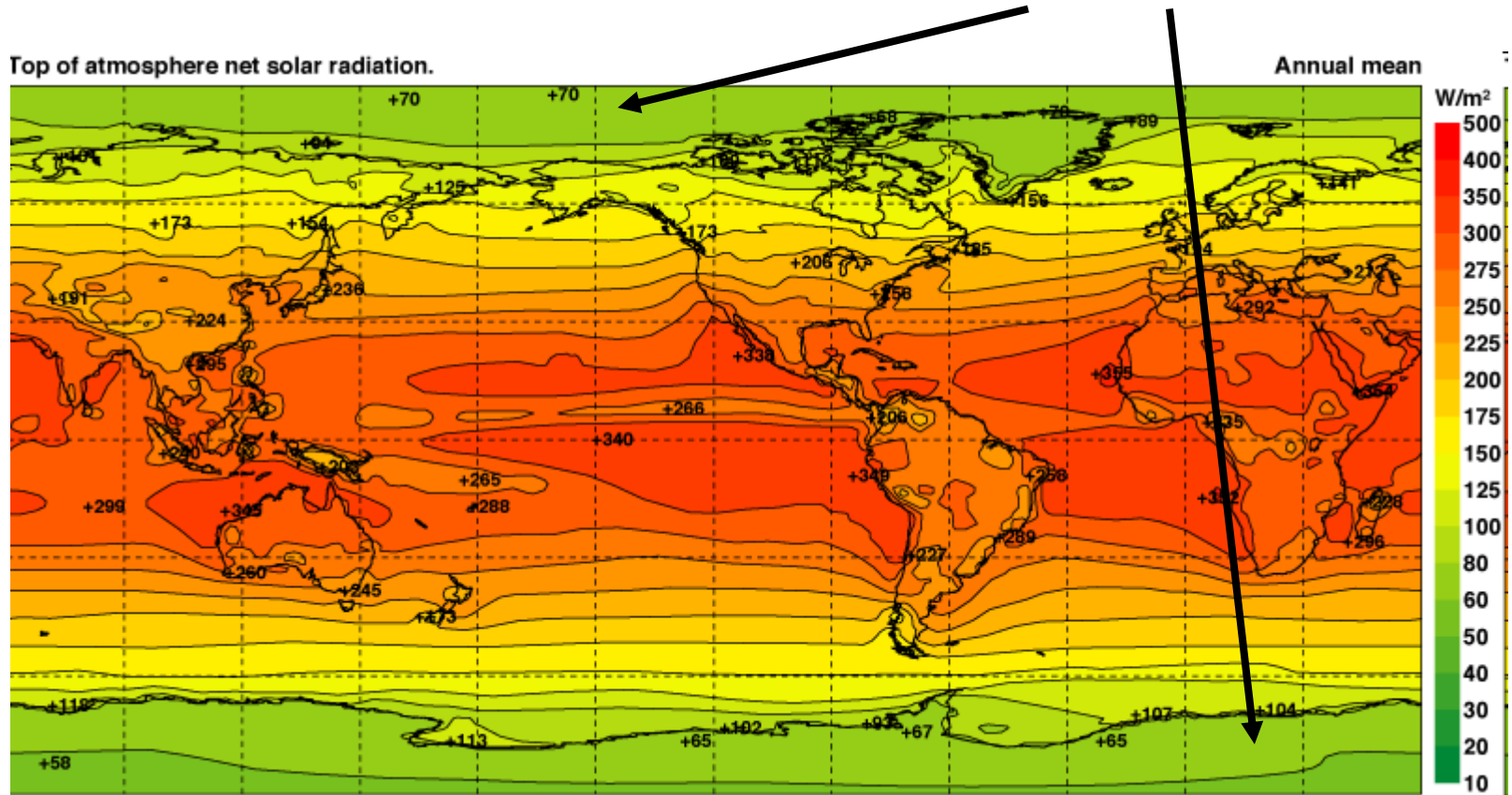
Annual mean



ECMWF, ERA40



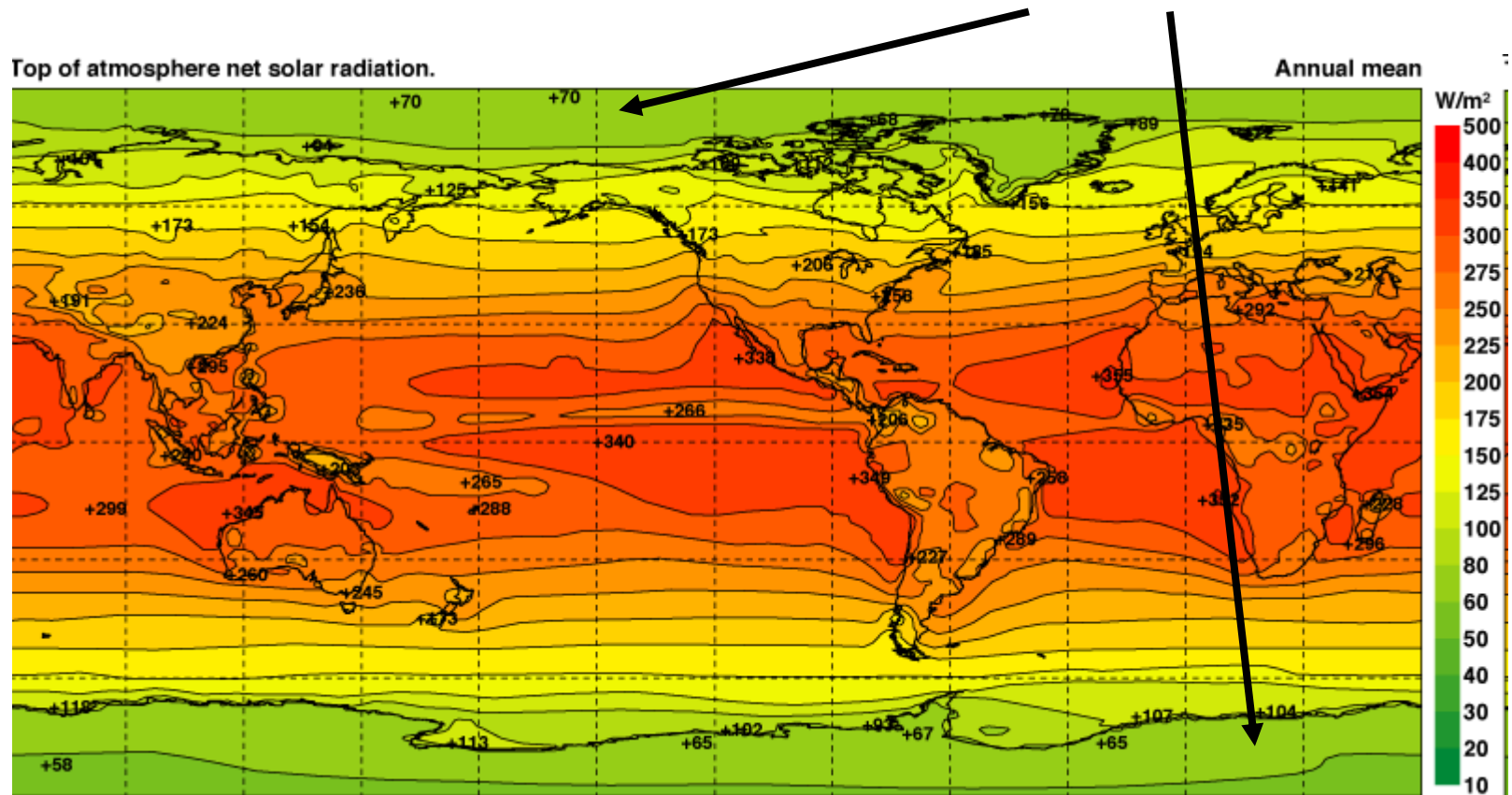
Average solar zenith angle large in polar regions



ECMWF, ERA40



High albedo due to ... low cloud and snow



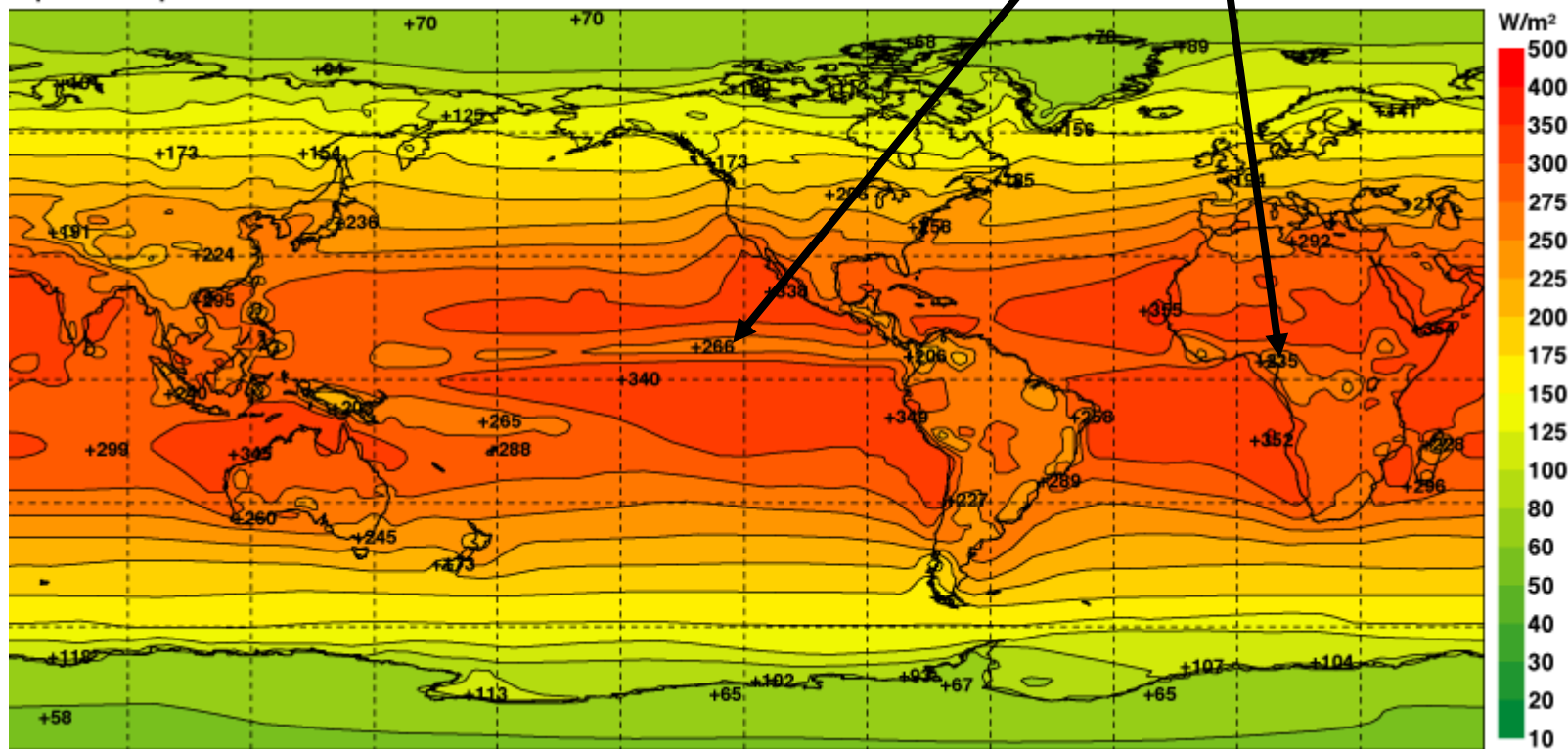
ECMWF, ERA40



High albedo due to ... persistent high clouds

Top of atmosphere net solar radiation.

Annual mean



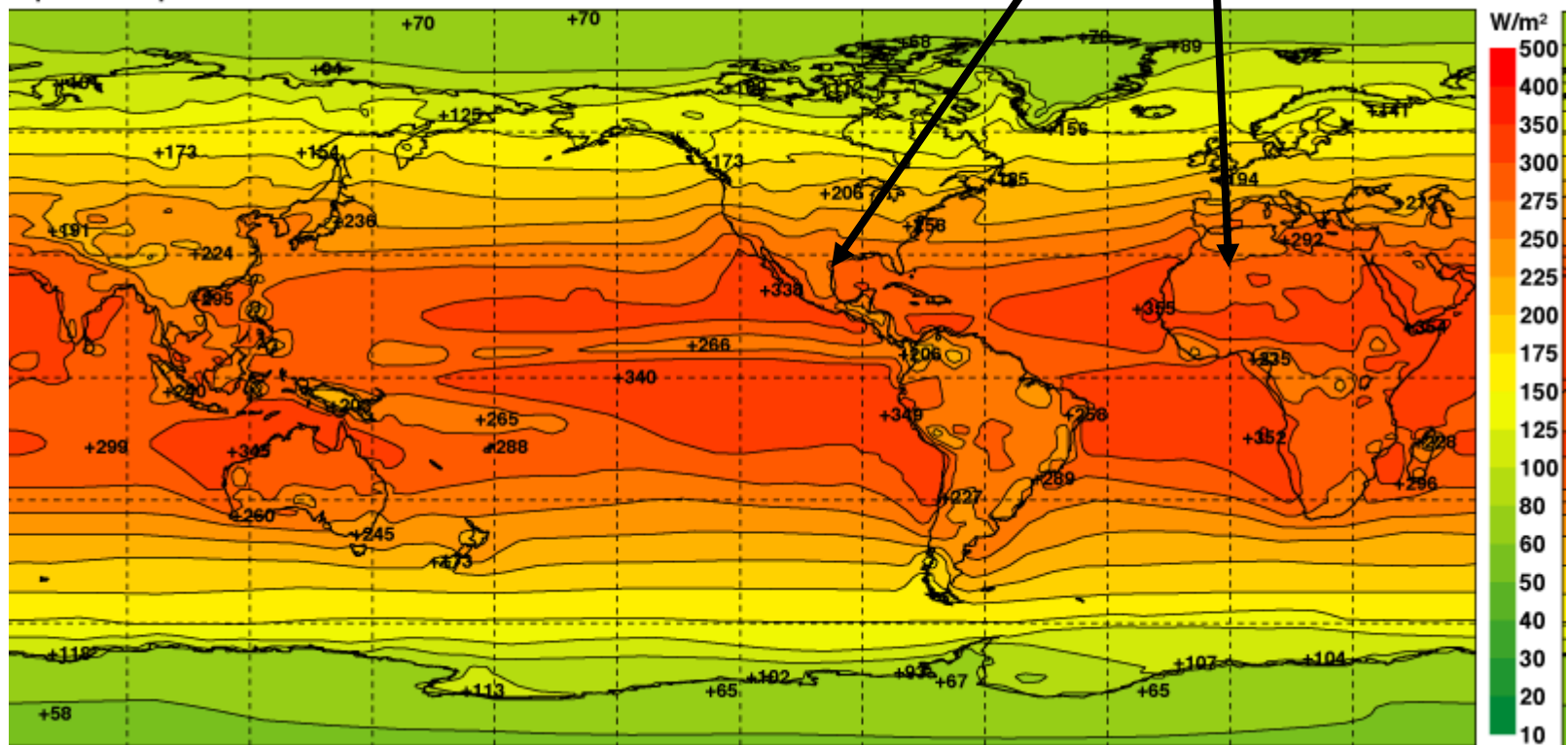
ECMWF, ERA40



High albedo due to ... bright surfaces

Top of atmosphere net solar radiation.

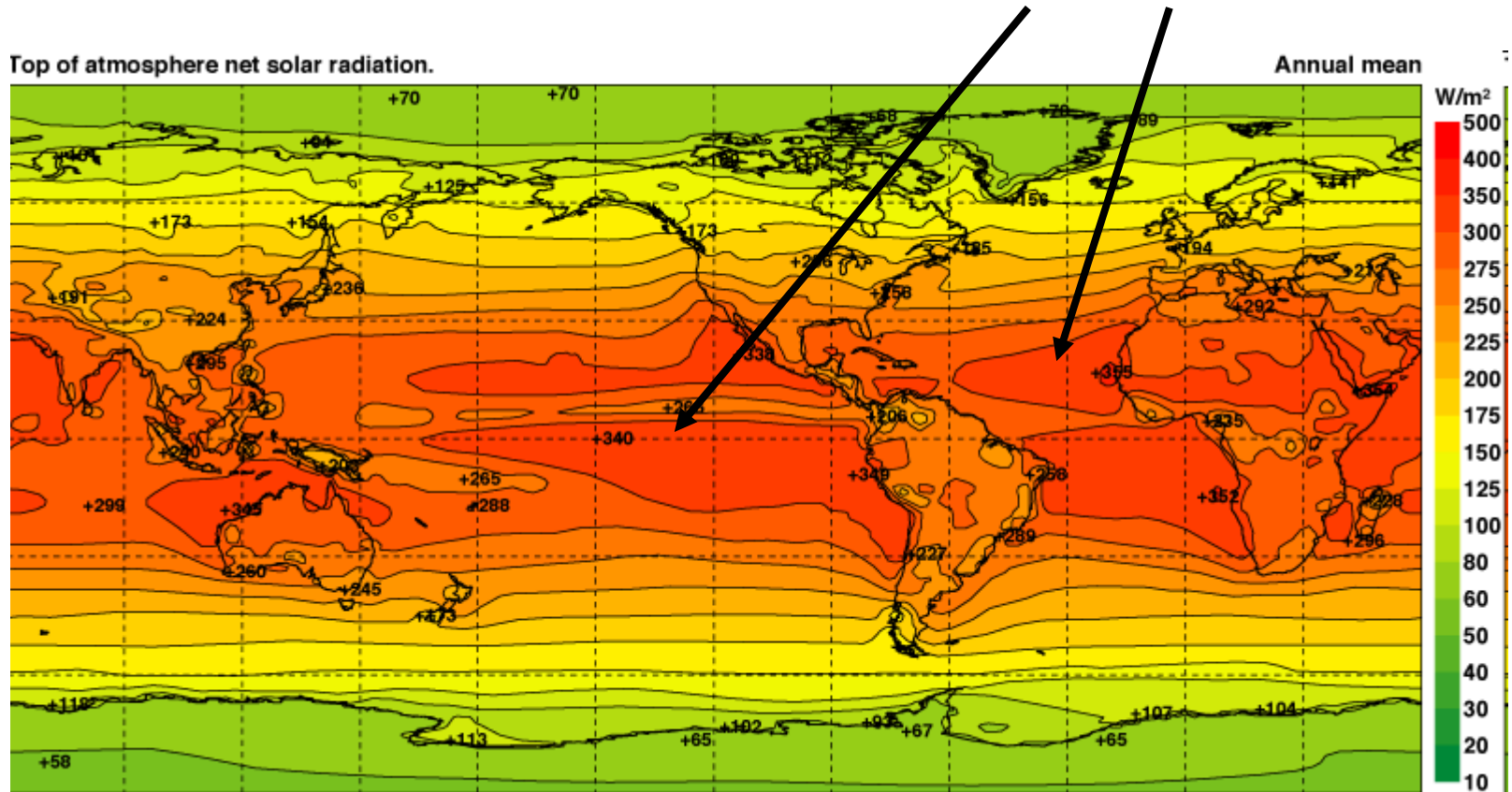
Annual mean



ECMWF, ERA40



Intrinsically low albedo of ocean surfaces



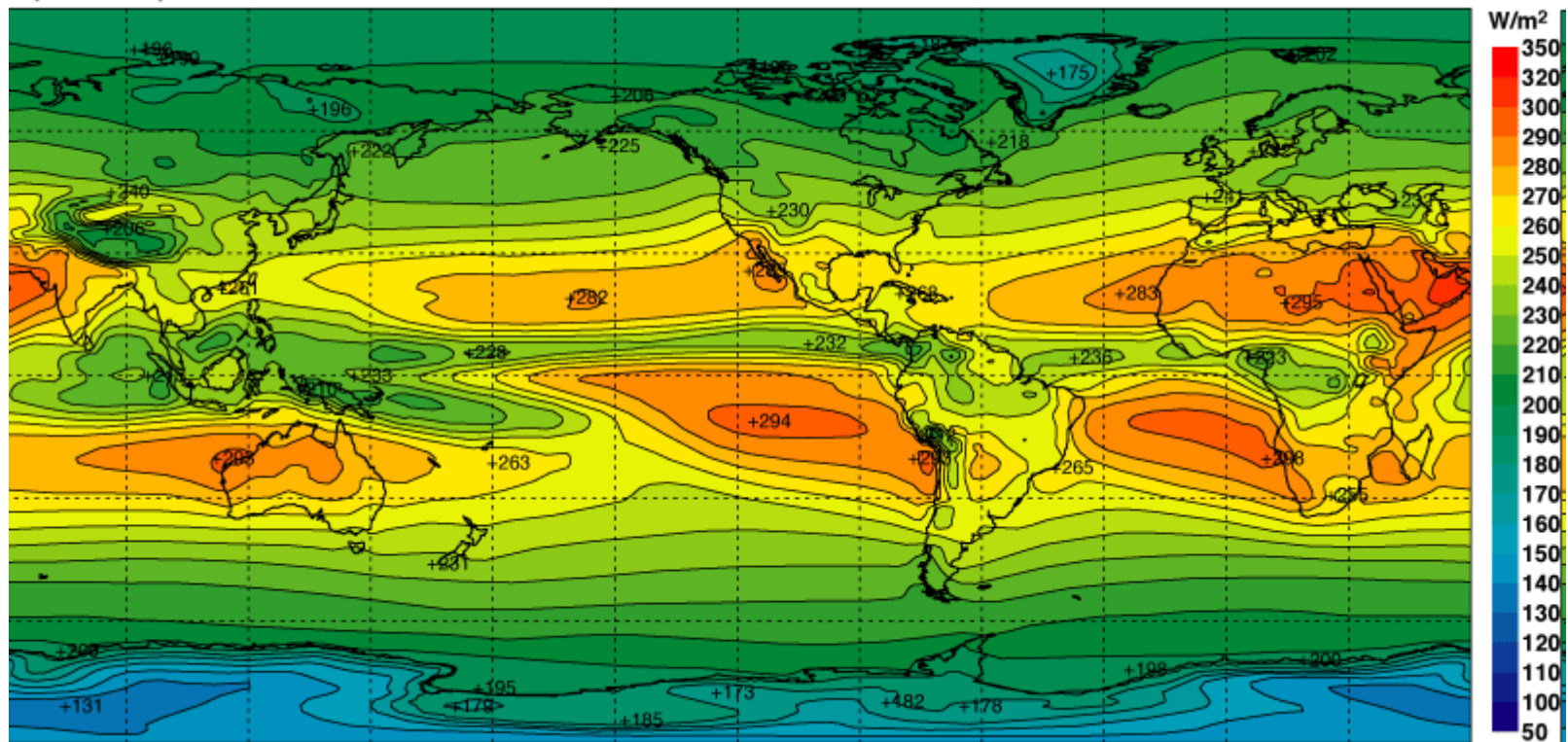
ECMWF, ERA40



Outgoing longwave radiation (emitted)

Top of atmosphere net thermal radiation

Annual mean



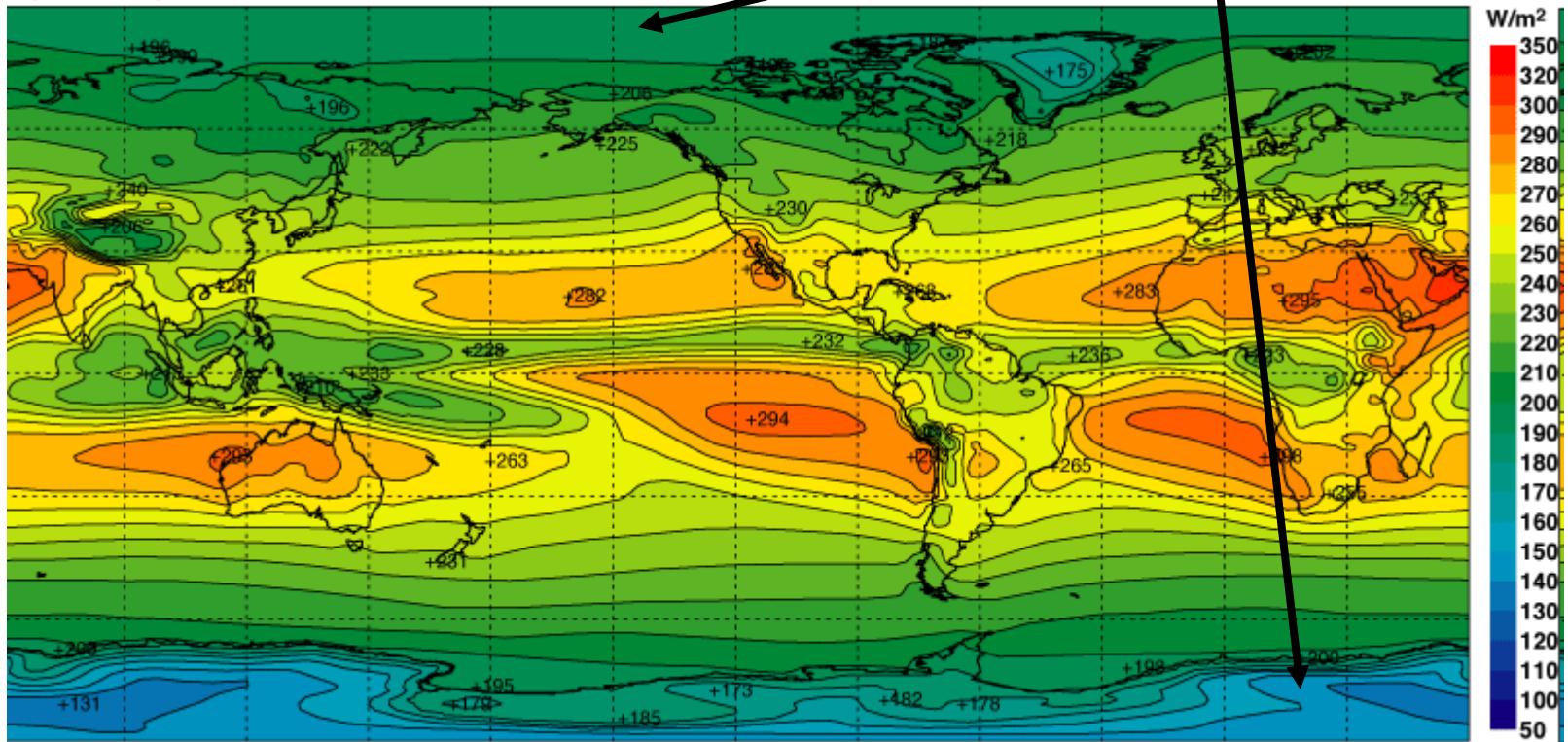
ECMWF, ERA40



Cold surface and atmosphere

Top of atmosphere net thermal radiation

Annual mean



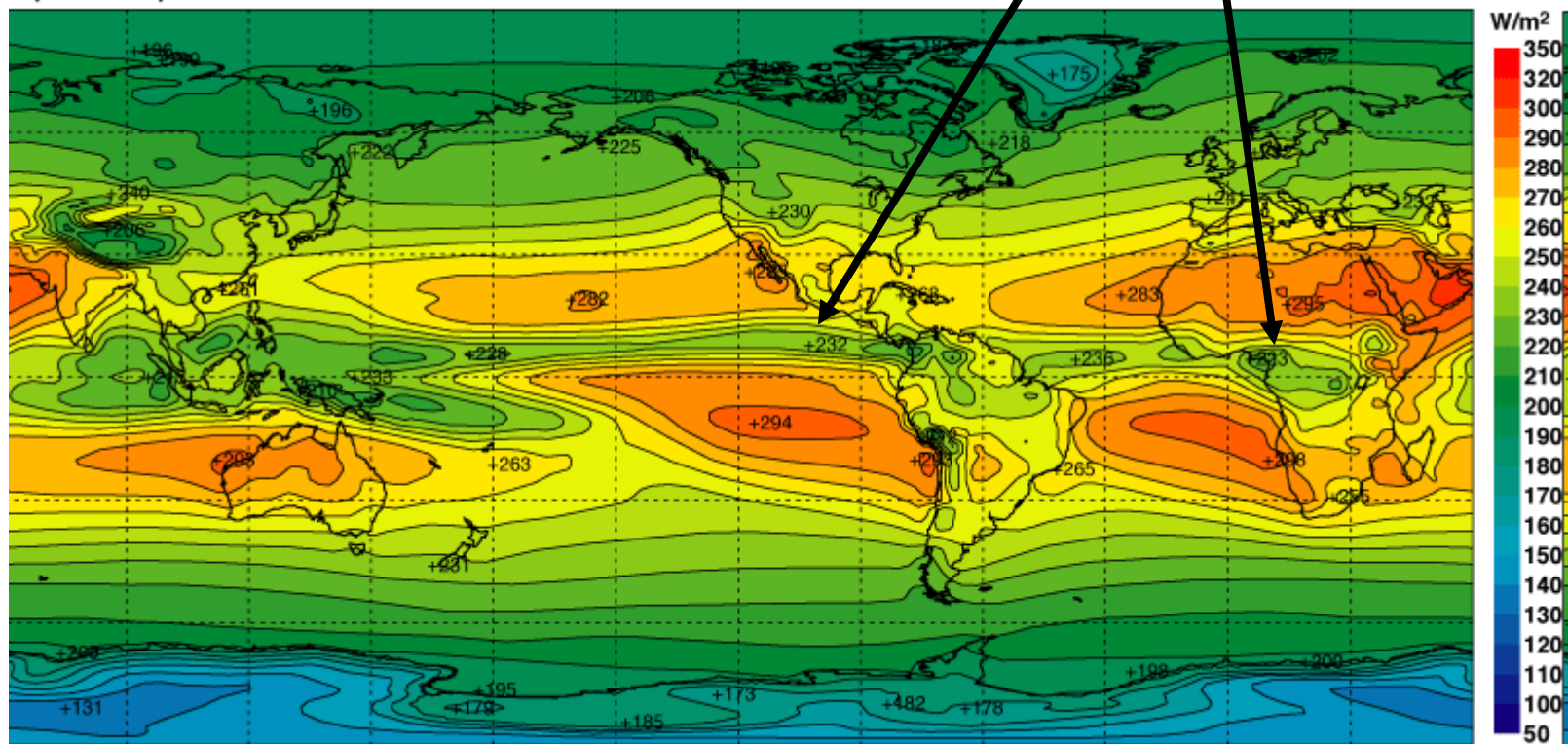
ECMWF, ERA40



Persistent high level clouds

Top of atmosphere net thermal radiation

Annual mean



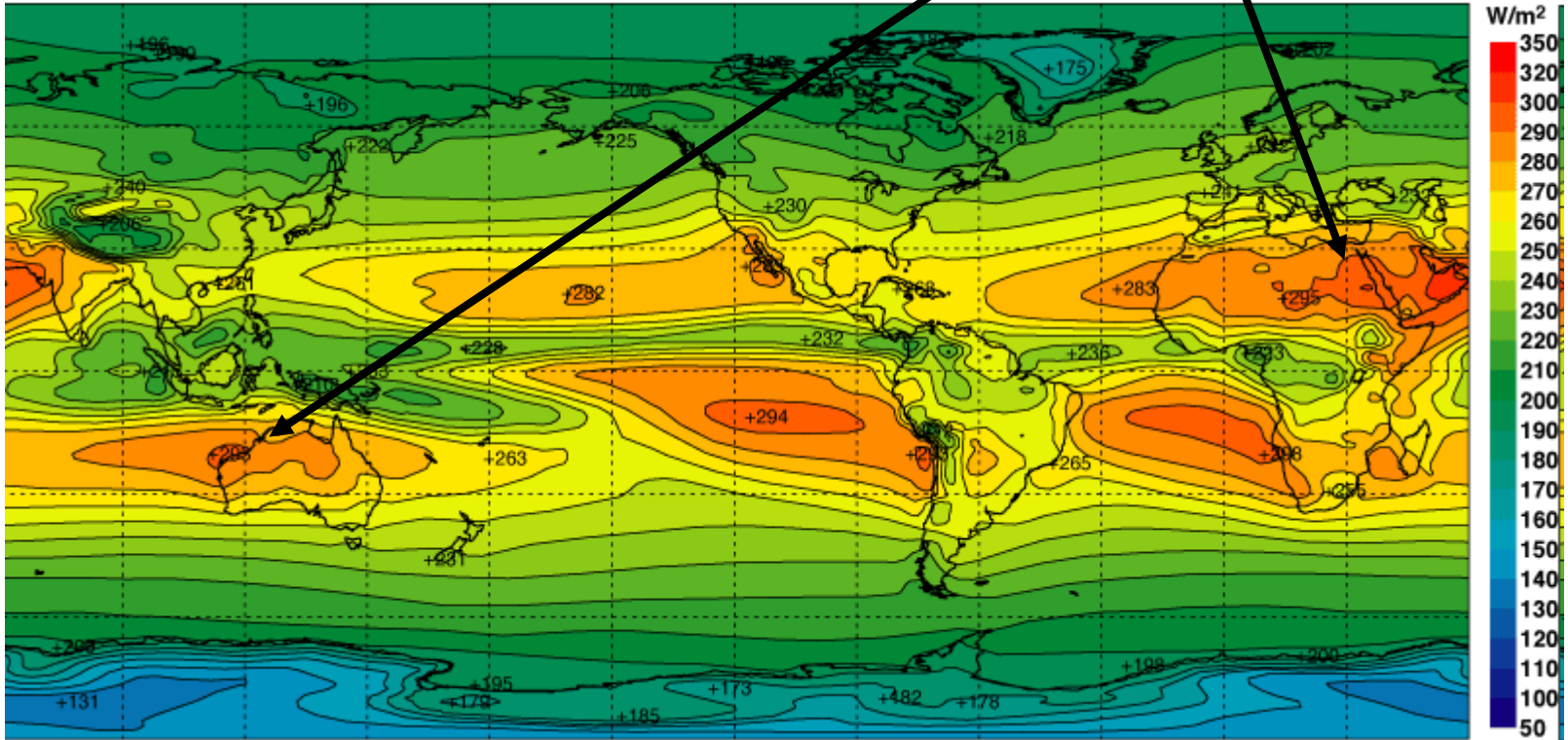
ECMWF, ERA40



Warm surface, dry atmosphere

Top of atmosphere net thermal radiation

Annual mean



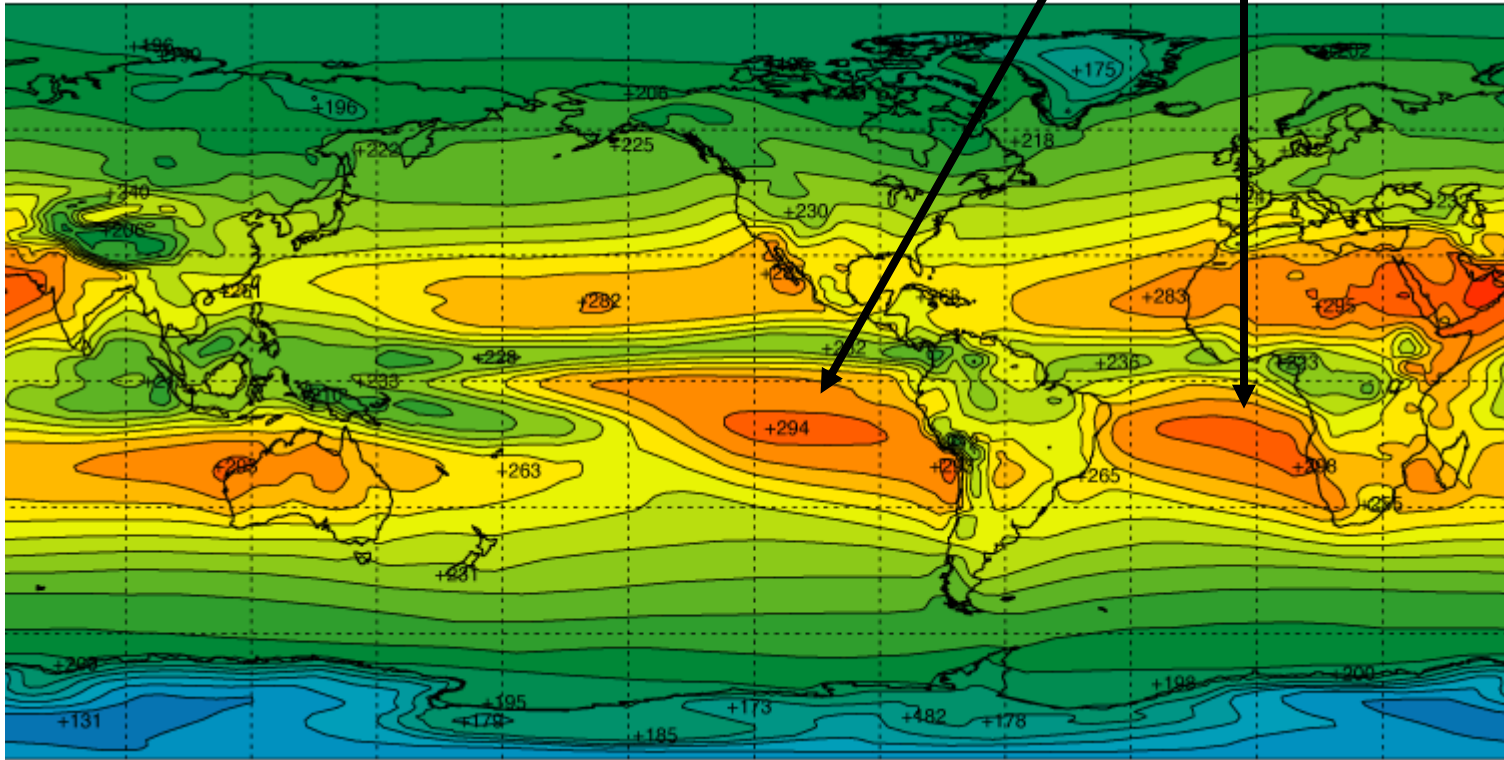
ECMWF, ERA40



Tropical oceans, low cloudiness

Top of atmosphere net thermal radiation

Annual mean



ECMWF, ERA40

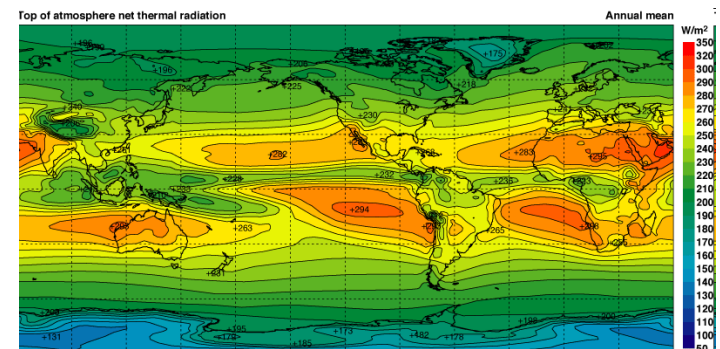
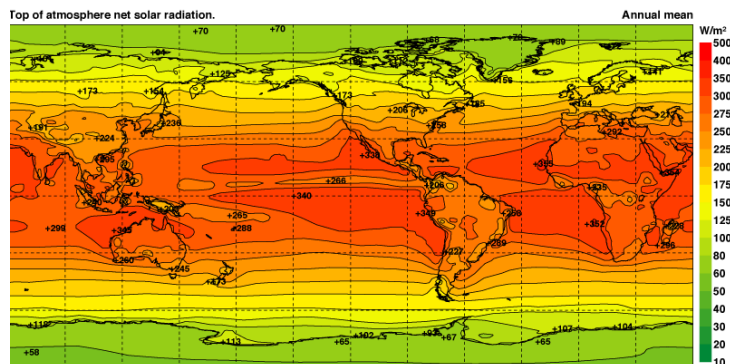


Net incoming radiation

Absorbed solar

... minus ...

emitted longwave



... and take average over latitude circles



Net incoming radiation

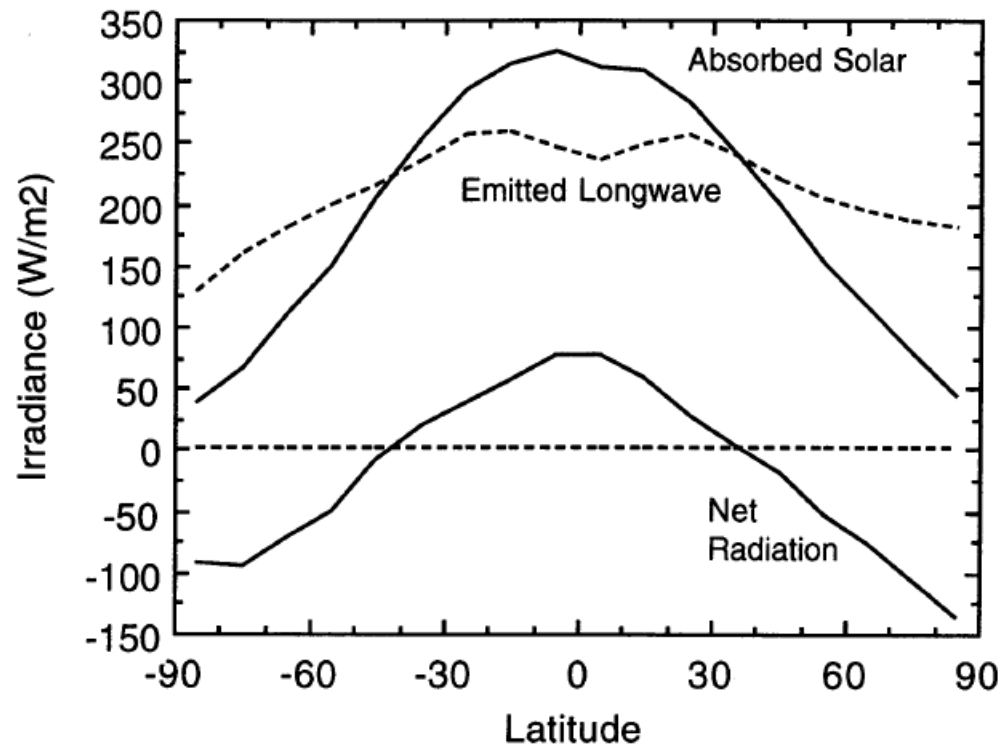


Fig. 2.12 Graphs of annual-mean absorbed solar radiation, OLR, and net radiation averaged around latitude circles.

Hartmann, 1994



Net incoming radiation

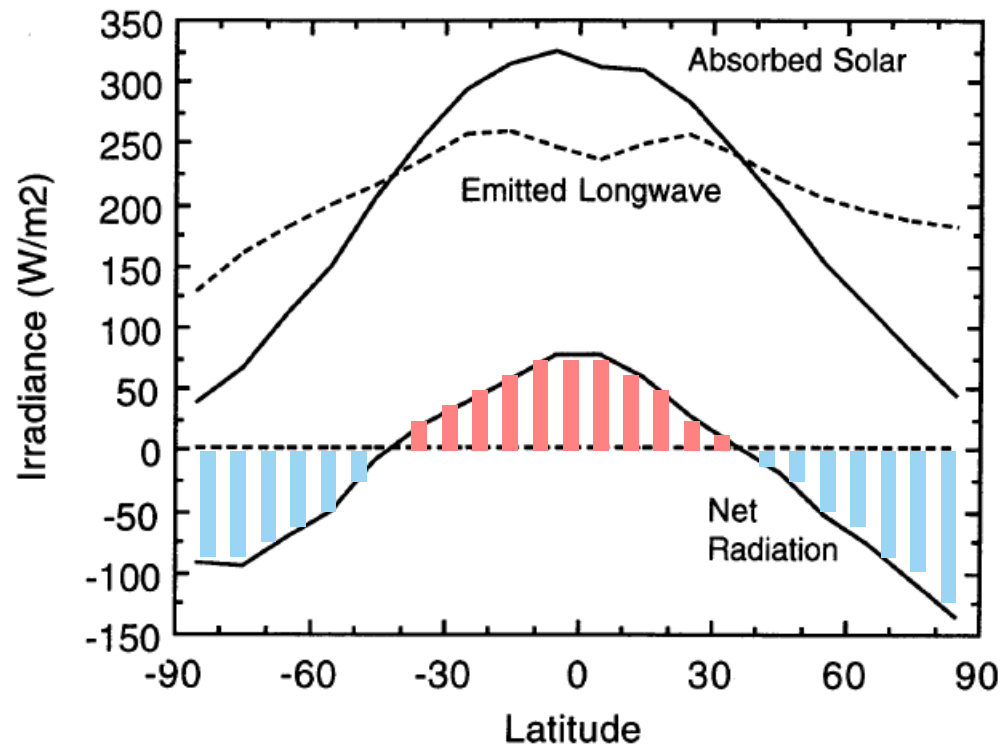


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Hartmann, 1994



... Poleward energy flux

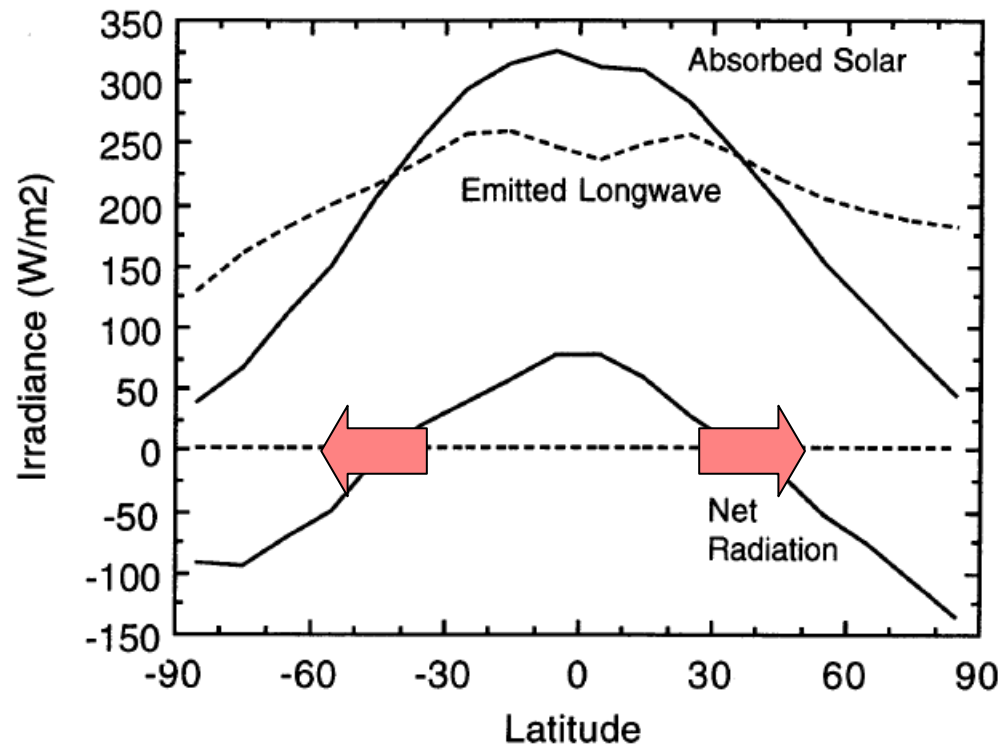


Fig. 2.12 Graphs of annual-mean absorbed solar radiation, OLR, and net radiation averaged around latitude circles.

Hartmann, 1994



Upper boundary condition

- **Conclusion**
 - Purely radiative
 - Relatively straightforward



2. Interaction with the underlying surface

- **Lower boundary condition of the atmosphere**
 - Physical balances of energy, momentum and water
 - Biogeochemical balances of
 - gases and vapours (CO₂, VOCs, ...)
 - particulate matter (aerosol)
 - Interactions with other Earth system components
 - Highly heterogenous surface properties
- **Next: surface balances of energy and momentum**



2. Surface energy balance

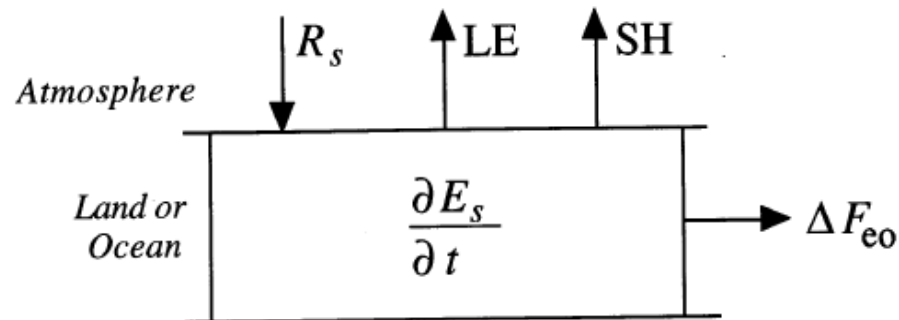


Fig. 4.1 Diagram showing the relationship of the various terms in the surface energy balance (R_s = net radiation, LE = evaporative cooling, SH = sensible cooling, $\partial E_s / \partial t$ = heat storage below the surface, ΔF_{eo} = divergence of horizontal energy flux below the surface).

Hartmann, 1994



2. Surface energy balance

Net radiative flux ↔ surface albedo variations

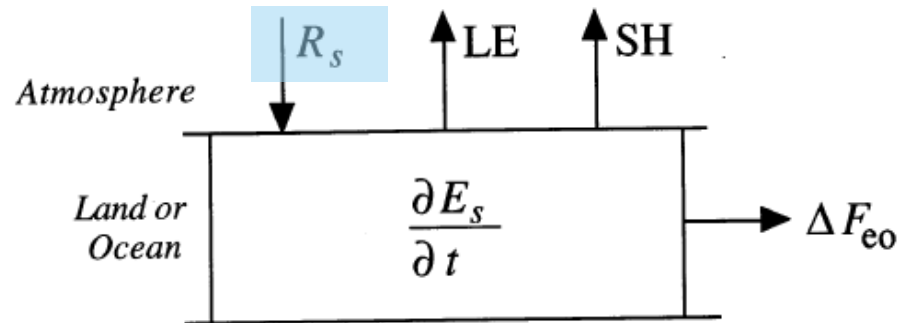


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Hartmann, 1994



2. Surface energy balance

Latent heat flux ↔ hydrological cycle

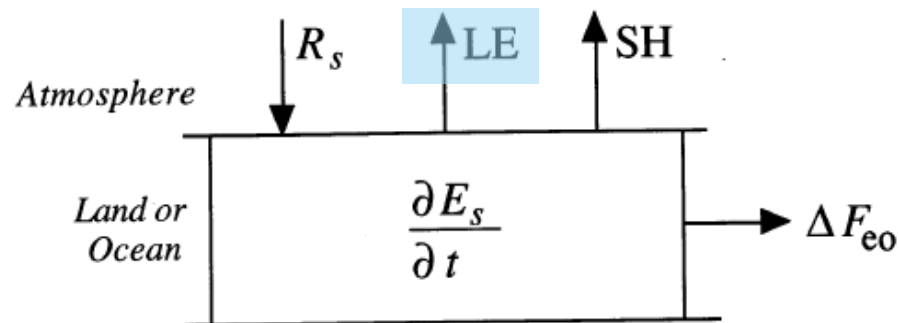


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Hartmann, 1994



2. Surface energy balance

Net radiative flux ↔

↔ hydrological cycle

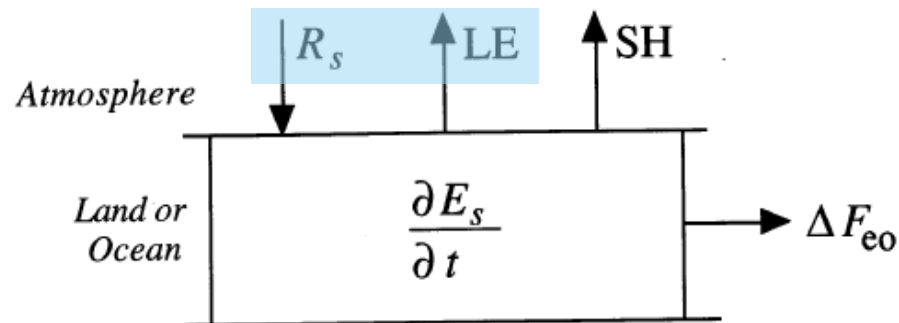


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Hartmann, 1994



2. Surface energy balance

- **Conclusion**

- important for understanding the climate
- far more complicated than the energy balance at the upper boundary



Surface momentum balance



Surface momentum balance

- **Angular momentum of the Earth system is conserved**

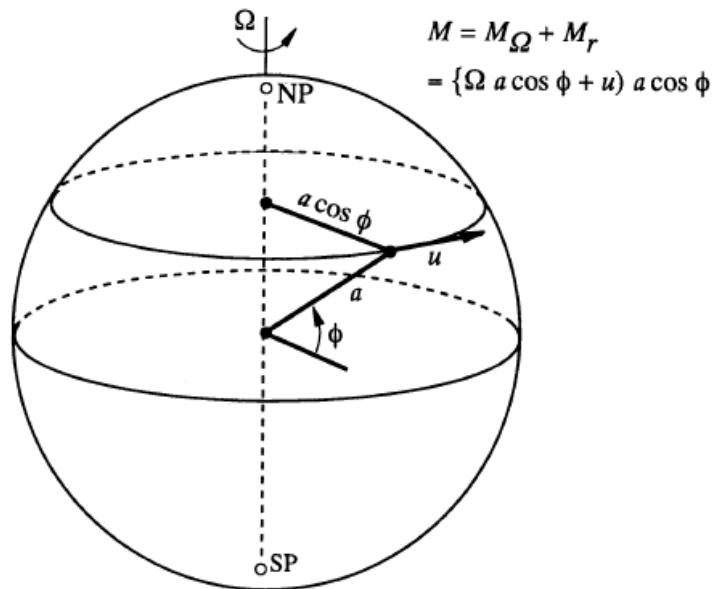


Fig. 6.14 Diagram showing the component of angular momentum about the axis of rotation of Earth. [From Peixóto and Oort (1984). Reprinted with permission from the American Physical Society.]





Conservation of angular momentum

- **No external torque**
- **Transfer of angular momentum to/from atmosphere via frictional forces and pressure forces acting on mountains**
- **Boundary layer processes are important**



Conservation of angular momentum

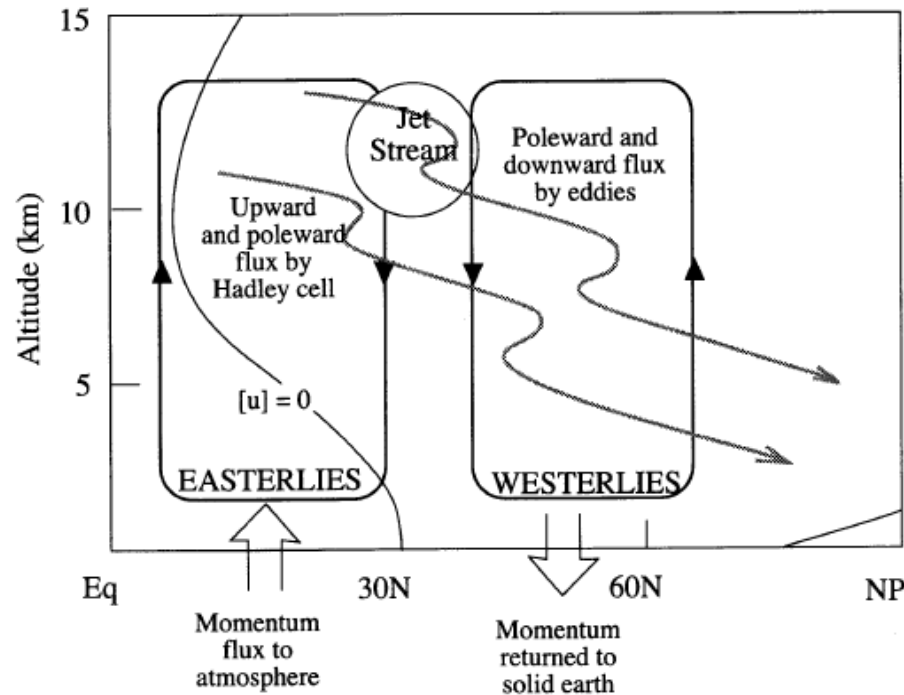


Fig. 6.16 Schematic illustration of the flow of angular momentum from Earth through the atmosphere and back to Earth.

Hartmann, 1994



Conservation of angular momentum

- **Poses a strong constraint for the atmospheric general circulation**

(For instance: westerly winds everywhere would drain the angular momentum of the atmosphere very soon)

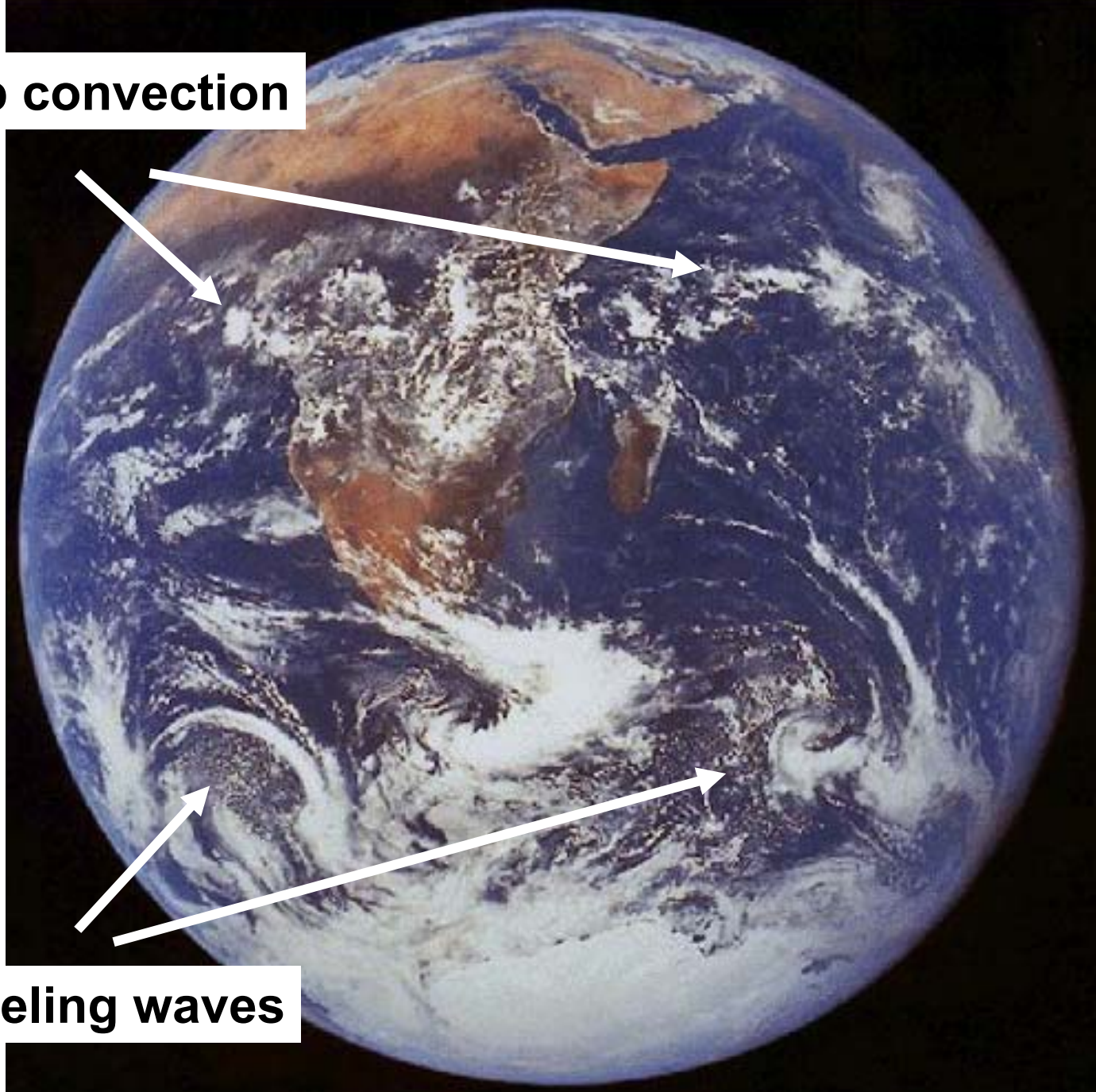


3. Internal dynamics

Lee vortices downstream of Jan Mayen Island (NASA)



Tropical deep convection



Midlatitude traveling waves



Governing laws of atmospheric motions

Momentum conservation

$$\frac{\partial U}{\partial t} + \frac{1}{a \cos^2 \theta} \left\{ U \frac{\partial U}{\partial \lambda} + V \cos \theta \frac{\partial U}{\partial \theta} \right\} + \dot{\eta} \frac{\partial U}{\partial \eta} = fV - \frac{1}{a} \left\{ \frac{\partial \phi}{\partial \lambda} + RT_v \frac{\partial}{\partial \lambda} \ln p \right\} + P_u + K_u$$

$$\frac{\partial V}{\partial t} + \frac{1}{a \cos^2 \theta} \left\{ U \frac{\partial V}{\partial \lambda} + V \cos \theta \frac{\partial V}{\partial \theta} + \sin \theta (U^2 + V^2) \right\} + \dot{\eta} \frac{\partial V}{\partial \eta} = -fV - \frac{\cos \theta}{a} \left\{ \frac{\partial \phi}{\partial \theta} + RT_v \frac{\partial}{\partial \theta} \ln p \right\} + P_v + K_v$$

Energy conservation

$$\frac{\partial T}{\partial t} + \frac{1}{a \cos^2 \theta} \left\{ U \frac{\partial T}{\partial \lambda} + V \cos \theta \frac{\partial T}{\partial \theta} \right\} + \dot{\eta} \frac{\partial T}{\partial \eta} = \frac{\kappa T_v \omega}{p} + P_T + K_T$$

Mass conservation

$$\frac{\partial}{\partial t} (\ln p_s) = -\frac{1}{p_s} \int_0^1 \nabla \cdot \left(\mathbf{v}_h \frac{\partial p}{\partial \eta} \right) d\eta$$

Ideal gas law

$$p = \rho RT$$

Moisture equation

$$\frac{\partial q}{\partial t} + \frac{1}{a \cos^2 \theta} \left\{ U \frac{\partial q}{\partial \lambda} + V \cos \theta \frac{\partial q}{\partial \theta} \right\} + \dot{\eta} \frac{\partial q}{\partial \eta} = P_q + K_q$$



Governing laws of atmospheric motions

Momentum conservation

$$\frac{\partial U}{\partial t} + \frac{1}{a \cos^2 \theta} \left\{ U \frac{\partial U}{\partial \lambda} + V \cos \theta \frac{\partial U}{\partial \theta} \right\} + \dot{\eta} \frac{\partial U}{\partial \eta} = fV - \frac{1}{a} \left\{ \frac{\partial \phi}{\partial \lambda} + RT_v \frac{\partial}{\partial \lambda} \ln p \right\} + P_u + K_u$$

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Coupled system
of

Partial differential equations



Governing laws of atmospheric motions

Momentum conservation

$$\frac{\partial U}{\partial t} + \frac{1}{a \cos^2 \theta} \left\{ U \frac{\partial U}{\partial \lambda} + V \cos \theta \frac{\partial U}{\partial \theta} \right\} + \dot{\eta} \frac{\partial U}{\partial \eta} = fV - \frac{1}{a} \left\{ \frac{\partial \phi}{\partial \lambda} + RT_v \frac{\partial}{\partial \lambda} \ln p \right\} + P_u + K_u$$

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Energy conservation

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Mass conservation

$$\frac{\partial}{\partial t} (\ln p_s) = -\frac{1}{p_s} \int_0^1 \nabla \cdot \left(\mathbf{v}_h \frac{\partial p}{\partial \eta} \right) d\eta$$

Non-linear

Ideal gas law

$$p = \rho RT$$

Moisture equation

$$\frac{\partial q}{\partial t} + \frac{1}{a \cos^2 \theta} \left\{ U \frac{\partial q}{\partial \lambda} + V \cos \theta \frac{\partial q}{\partial \theta} \right\} + \dot{\eta} \frac{\partial q}{\partial \eta} = P_q + K_q$$



Governing laws of atmospheric motions

Momentum conservation

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Energy conservation

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Mass conservation

$$\frac{\partial}{\partial t} (\ln p_s) = -\frac{1}{p_s} \int_0^1 \nabla \cdot \left(\mathbf{v}_h \frac{\partial p}{\partial \eta} \right) d\eta$$

Dissipative

Ideal gas law

$$p = \rho RT$$

Moisture equation

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Governing laws of atmospheric motions

Momentum conservation

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Forced

Ideal gas law

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Moisture equation

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Governing laws of atmospheric motions

Momentum conservation

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Ideal gas law

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Moisture equation

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Chaotic
=

Deterministic non-periodic flow



... Poleward energy flux

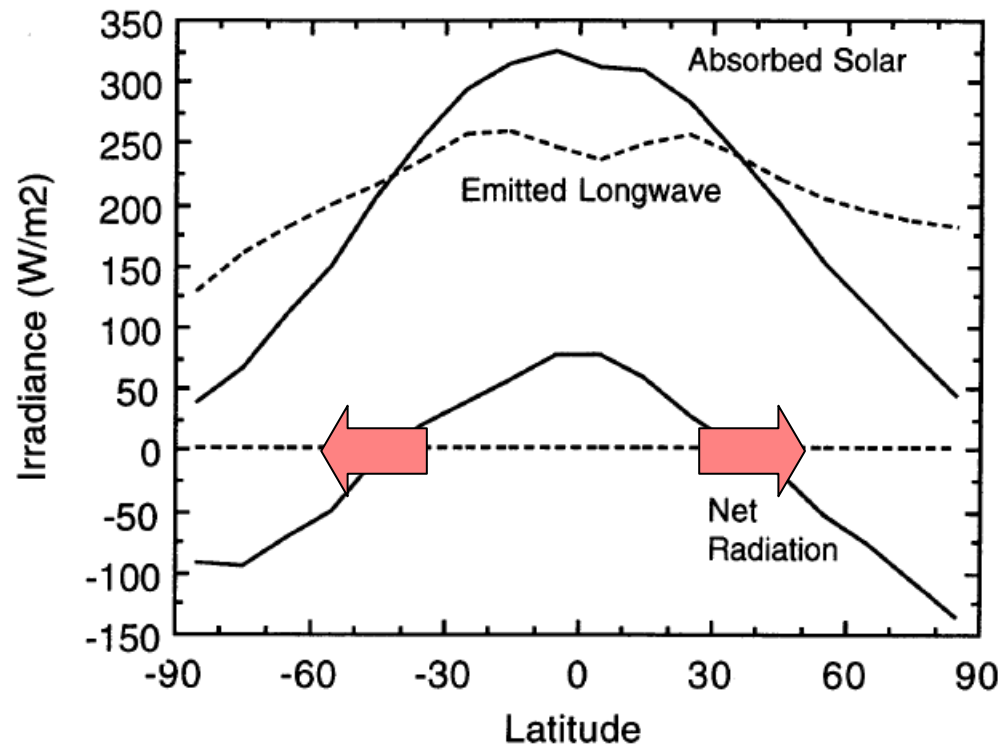


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Hartmann, 1994



... Poleward energy flux

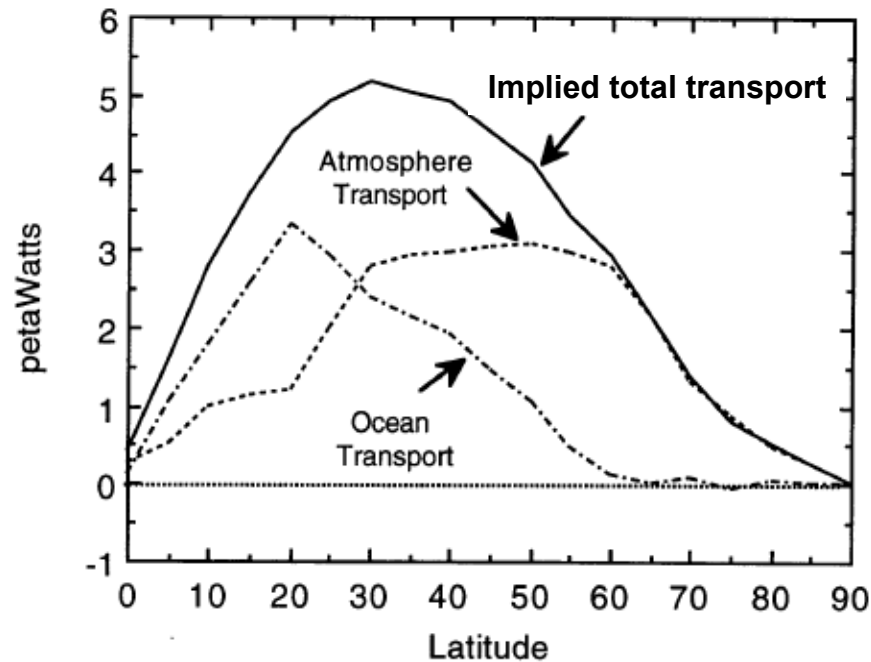


Fig. 2.14 Meridional transport of energy for annual-mean conditions. Net radiation and atmospheric transport are estimated from observations; ocean transport is calculated as a residual in the energy balance. [Adapted from Vonder Haar and Oort (1973). Used with permission from the American Meteorological Society.]

Hartmann, 1994



... Poleward energy flux

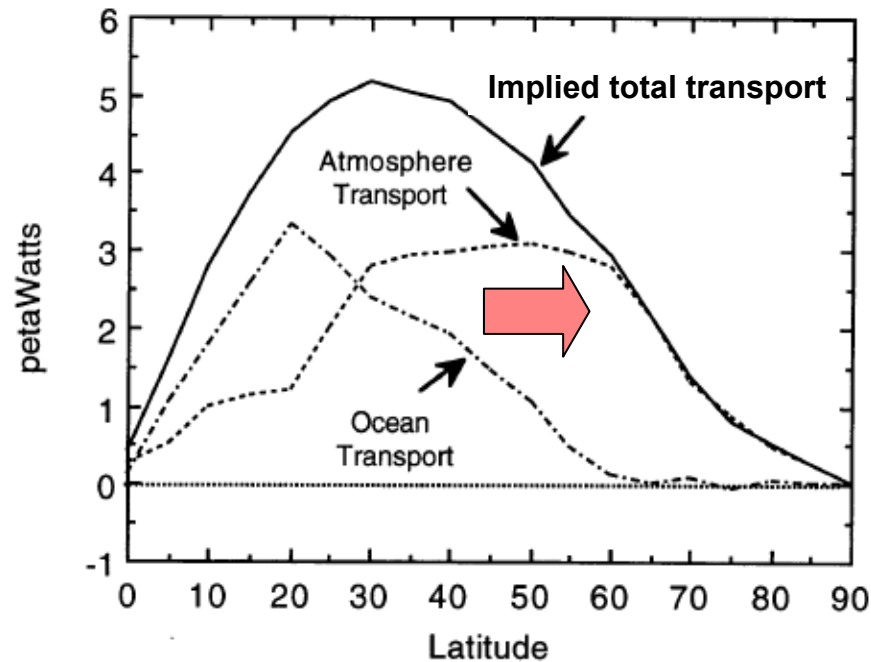
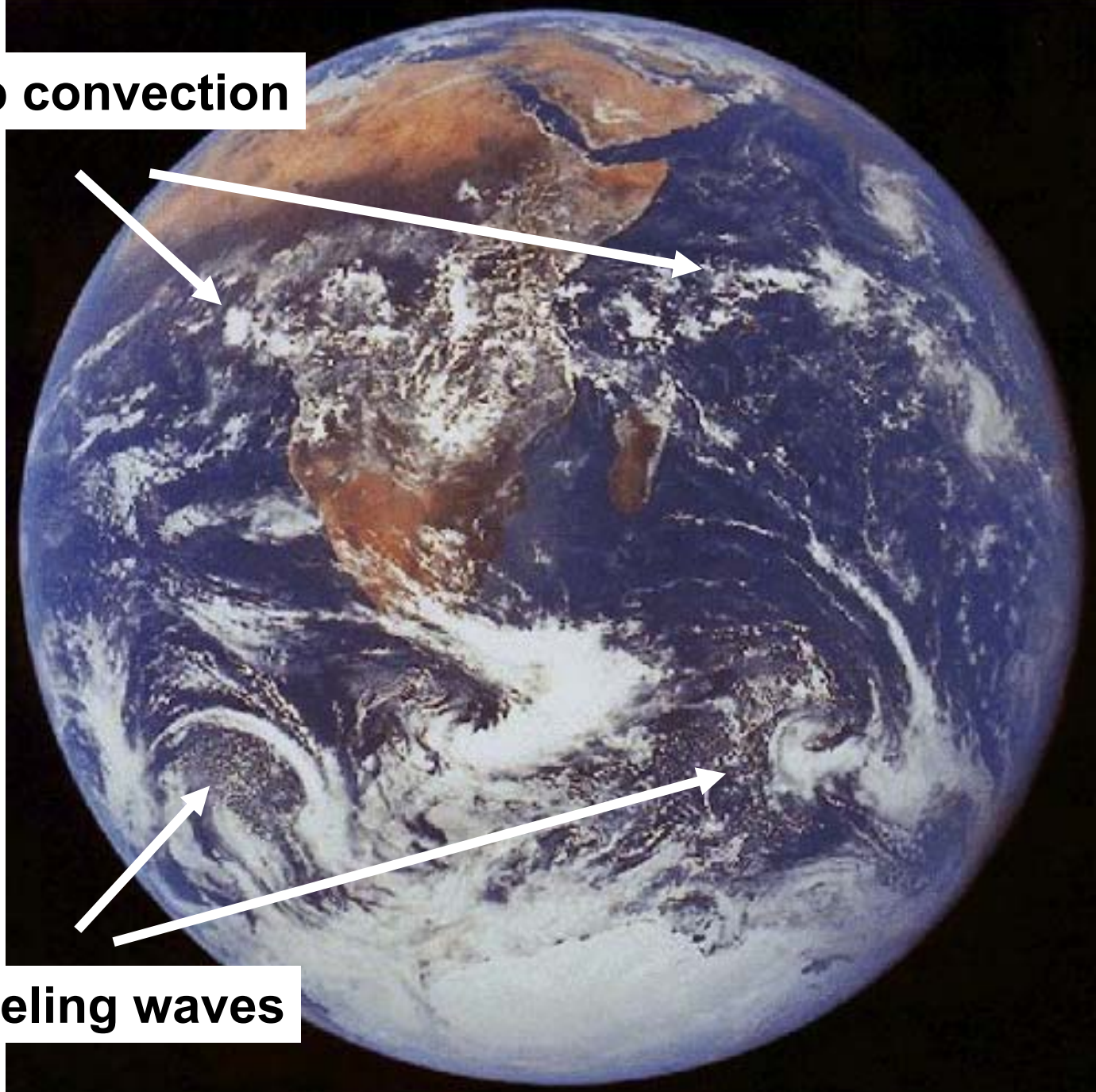


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Hartmann, 1994

Tropical deep convection



Midlatitude traveling waves



... Poleward energy flux

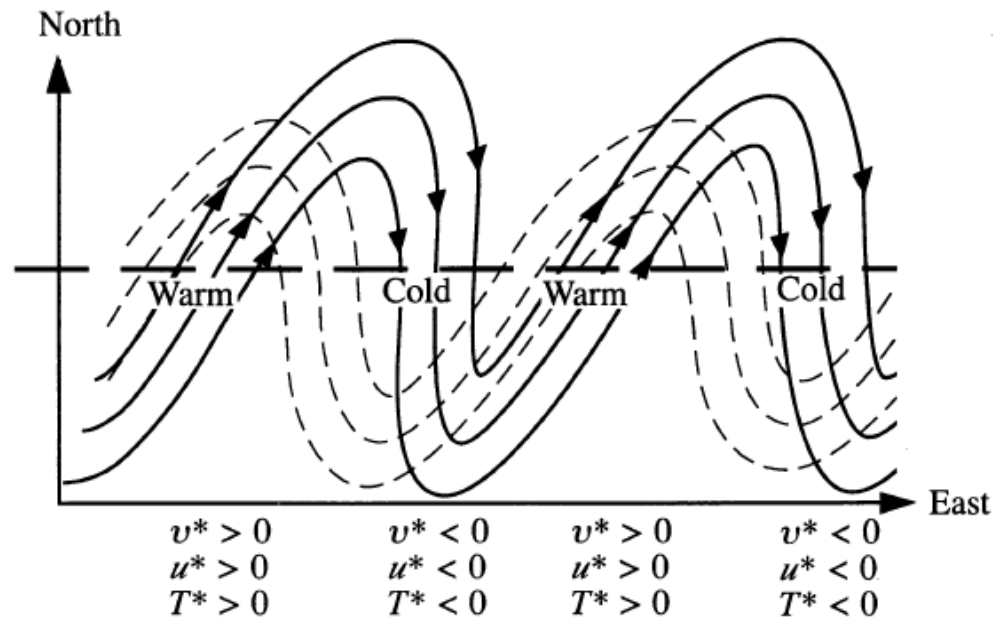


Fig. 6.6 Schematic of the streamlines (solid) and isotherms (dashed) associated with a large-scale atmospheric disturbance in midlatitudes of the Northern Hemisphere. Arrows along the streamline contour indicate the direction of wind velocity. The streamlines correspond approximately to lines of constant pressure, since the winds are nearly geostrophic. The signs of the deviations of the wind components from their zonal-average values are shown to illustrate that the NE–SW tilt of the streamlines indicates a northward zonal momentum transport, and the westward phase shift of the temperature wave relative to the pressure wave gives a northward heat transport.

Hartmann, 1994



... Poleward energy flux

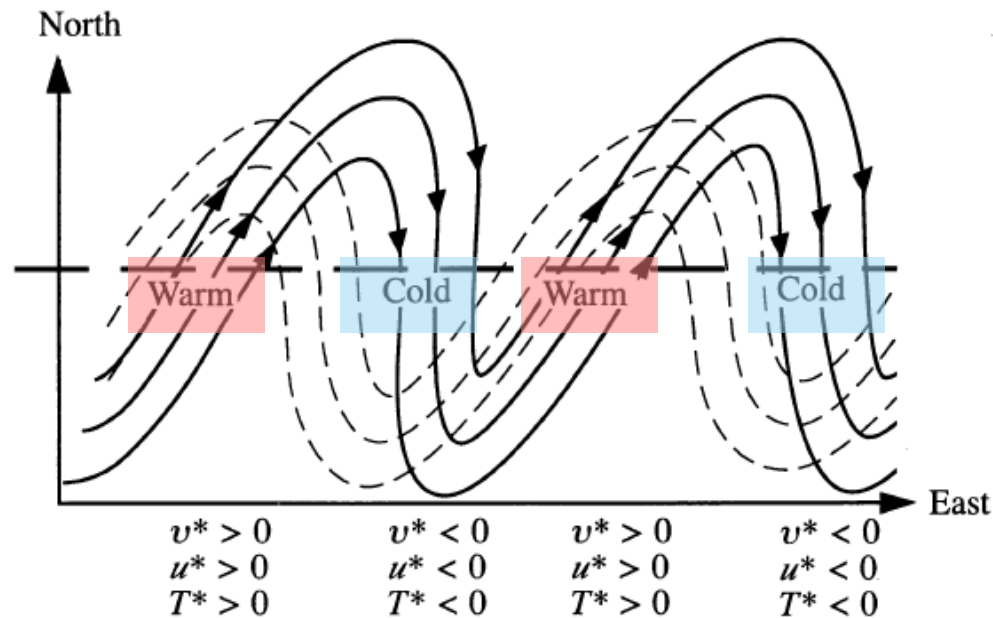


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Hartmann, 1994



... Poleward momentum flux

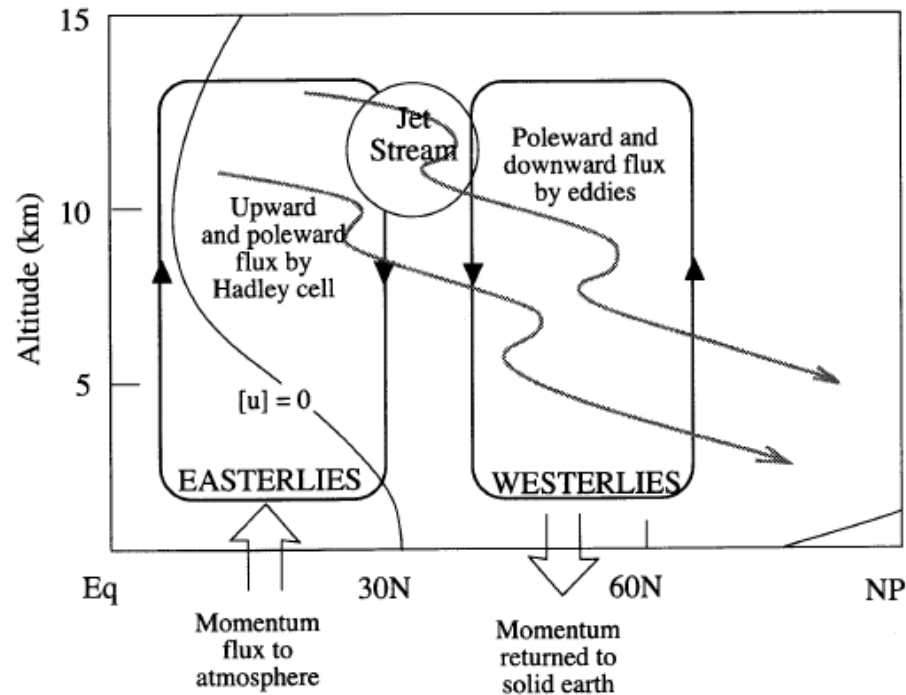


Fig. 6.16 Schematic illustration of the flow of angular momentum from Earth through the atmosphere and back to Earth.



Poleward fluxes of energy and momentum

- **Mid-latitude traveling waves ("depressions")**
- **Especially in the Southern Hemisphere where stationary waves are weak**



4. Concluding remarks

1. Interaction with the surrounding Universe

- Upper boundary condition for the Earth system

2. Interaction with the underlying surface

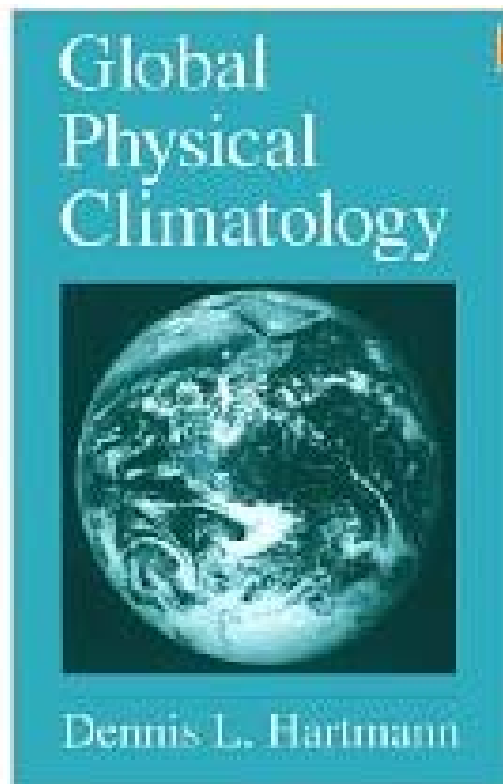
- Lower boundary condition for the atmosphere
- Non-linear interaction with other Earth system components

3. Internal dynamics

→ Approaches for Earth system model validation



Reference





Many thanks!

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